

**Guidelines**

- (i) You can reuse without demonstration all the results introduced in the course (by only citing the considered slide number for instance);
- (ii) An essential part of the work is to provide valuable comments of your results.

**Temperature fluctuations in buoyancy-driven flows**

The turbulent kinetic energy budget, which has been established during the lectures, is here revisited to include temperature effects in order to illustrate the possible competition between kinetic and thermal turbulence. The fluid is a perfect gas and the flow is assumed to be incompressible,  $\nabla \cdot \mathbf{u} = 0$ . The Navier-Stokes equation and the energy conservation written for the temperature  $T$ , are recalled below in a conservative form

$$\begin{cases} \frac{\partial(\rho \mathbf{u})}{\partial t} + \nabla \cdot (\rho \mathbf{u} \mathbf{u}) = -\nabla p + \nabla \cdot \overline{\boldsymbol{\tau}} + \rho \mathbf{g} \\ \frac{\partial(\rho c_p T)}{\partial t} + \nabla \cdot (\rho c_p T \mathbf{u}) = -\nabla \cdot \mathbf{q} + \overline{\boldsymbol{\tau}} : \nabla \mathbf{u} + \frac{Dp}{Dt} \end{cases} \quad (1)$$

where  $c_p$  is the specific heat at constant pressure ( $dh = c_p dT$  for a perfect gas),  $\mathbf{q} = -\lambda \nabla T$  (Fourier's law),  $\overline{\boldsymbol{\tau}} : \nabla \mathbf{u} = \tau_{ij} \partial u_i / \partial x_j$  is the viscous dissipation, and  $g_i = -g \delta_{3i}$  is the gravity with  $g \simeq 9.81 \text{ m.s}^{-2}$ . All the fluid properties are assumed to be constant ( $\nu, c_p, \lambda$ ) to simplify algebra.

**A – Simplification of the equation of state.** In the framework of a low Mach number approximation, density  $\rho = \rho(T, p)$  can be considered as a function of temperature only, that is  $\rho \simeq \rho(T)$ , as shown below.

1. Derive the expressions of the thermal expansion coefficient  $\beta$  and of the isothermal compressibility coefficient  $\chi$  for a perfect gas. Calculate their numerical values for the reference state  $P_0 = 101325 \text{ Pa}$ ,  $T_0 = 288.15 \text{ K}$  and  $\rho_0 = 1.225 \text{ kg.m}^{-3}$ . As a reminder,

$$\beta = -\frac{1}{\rho} \left. \frac{\partial \rho}{\partial T} \right|_p \quad \chi = \frac{1}{\rho} \left. \frac{\partial \rho}{\partial p} \right|_T \quad (2)$$

2. By considering the Taylor series of density around its mean value  $\rho_0$ , show that

$$\frac{\Delta \rho}{\rho_0} = -\frac{\Delta T}{T_0} + \frac{\Delta p}{P_0} \quad (3)$$

where  $\Delta \rho = \rho - \rho_0$ ,  $\Delta T = T - T_0$  and  $\Delta p = p - P_0$  denote the small perturbation in density, temperature and pressure. Retrieve this result by considering the logarithmic differentiation of the perfect gas law.

3. Assuming that the total pressure of a fluid particle remains constant during its motion and using Eq. (3), demonstrate that

$$\left(1 + \frac{\gamma}{2} M^2\right) \frac{\Delta \rho}{\rho} \simeq -\frac{\Delta T}{T} - \gamma M^2 \frac{\Delta u}{U} \quad \text{with} \quad M = \frac{U}{c} \quad (4)$$

4. Plot the variation of  $\Delta \rho / \rho$  as a function of  $M$  for several values of  $\Delta T / T$  and  $\Delta u / U$ . Conclude this first part.

**B – Generalised Boussinesq approximation.** We now consider a reference steady state  $(\rho_0, P_0, \mathbf{u}_0 = 0, T_0)$  at equilibrium which satisfies the hydrostatic balance  $-\nabla P_0 + \rho_0 \mathbf{g} = 0$  and  $\nabla^2 T_0 = 0$ , and the presence of a turbulent atmospheric boundary layer. The flow variables are then decomposed as follows

$$\rho = \rho_0 + \rho' \quad p = P_0 + p^* = P_0 + \bar{P} + p' \quad \mathbf{u} = \bar{\mathbf{U}} + \mathbf{u}' \quad T = T_0 + T^* = T_0 + \bar{T} + \theta \quad (5)$$

to remove the effects of the static stratification. From the previous part A, the equation of state takes the form  $\rho \simeq \rho_0(1 - \beta T^*)$  with  $\beta = 1/T_0$ . The temperature  $T_0$  is almost constant for an adiabatic atmosphere at the scale of the turbulent field.

5. Show that the momentum conservation equation can be recast as follows

$$\rho_0(1 - \beta T^*) \frac{D\mathbf{u}}{Dt} = -\nabla p^* + \nabla \cdot \bar{\boldsymbol{\tau}} - \rho_0 \beta T^* \mathbf{g} \quad (6)$$

6. Deduce that for small temperature perturbations, the Boussinesq approximation is obtained : density fluctuations are taken into account in the buoyancy force  $\rho \mathbf{g}$  only.

**C – Temperature fluctuations in the lowest layer of Earth's atmosphere** – In this third part, we use the previous Boussinesq approximation to derive the turbulent kinetic energy budget including buoyancy effects. We assume a mean velocity field  $\bar{U}_1(x_3)$  consistent with geophysical practice.

7. Derive the equation for the mean temperature  $\bar{T}$  by introducing the Reynolds decomposition (5). How can we simplify this equation for low Mach number flows?

8. Show that the total mean heat flux  $\bar{q}^t$  can be written as

$$\bar{q}_j^t = -\rho_0 c_p \left( a \frac{\partial \bar{T}}{\partial x_j} - \overline{\theta u_j'} \right)$$

What does the coefficient  $a$  represent? Provide its dimension. How could you simplify the expression of the total heat flux for a high value of the Peclet number?

9. Compute the additional term associated with gravity in the equation for the fluctuating velocity component  $u_i'$ , and in the equation for the turbulent kinetic energy  $k_t$ .

10. In order to quantify the importance of the buoyancy force effects, we can compare the additional term in the transport equation for  $k_t$  with the production term  $\mathcal{P}$ , by introducing the Richardson number  $Ri$  defined as

$$Ri \equiv -\frac{\overline{\rho_0 g \theta u_3'} / T_0}{\mathcal{P}} \quad \text{where} \quad \mathcal{P} = -\rho_0 \overline{u_1' u_3'} \frac{\partial \bar{U}_1}{\partial x_3}$$

Show that if  $d\bar{T}/dx_3 < 0$ , the additional term acts as a production term for the turbulent kinetic energy. What is the sign of the Richardson number? Comment the dispersion of a pollutant for an unstable thermal stratification, that is when the temperature decreases with the altitude.

11. Repeat your analysis for the case of a stable thermal stratification.

12. A temperature inversion over the first few hundred meters is regularly observed in the atmospheric layer, leading to the trap of air pollution. Find a recent example of this phenomenon (internet, scientific paper written for a large audience), where the temperature inversion is documented and comment your finding in a few sentences.